



Designation: C 294 – 04

## Standard Descriptive Nomenclature for Constituents of Concrete Aggregates<sup>1</sup>

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*This standard has been approved for use by agencies of the Department of Defense.*

### 1. Scope

1.1 This descriptive nomenclature provides brief descriptions of some of the more commonly occurring, or more important, natural and artificial materials of which mineral aggregates are composed. The descriptions provide a basis for understanding these terms as applied to concrete aggregates. When appropriate, brief observations regarding the potential effects of using the natural and artificial materials in concrete are discussed.

NOTE 1—These descriptions characterize minerals and rocks as they occur in nature and blast-furnace slag or lightweight aggregates that are prepared by the alteration of the structure and composition of natural material. Information about lightweight aggregates are given in Specifications C 330, C 331, and C 332.

1.2 This standard does not include descriptions of constituents of aggregates used in radiation shielding concrete. See ASTM C 638.

### 2. Referenced Documents

#### 2.1 ASTM Standards:

- C 125 Terminology Relating to Concrete and Concrete Aggregates<sup>2</sup>
- C 227 Test Method for Potential Alkali Reactivity of Cement & Aggregate Combinations (Mortar Bar Method)<sup>2</sup>
- C 289 Test Method for Potential Reactivity of Aggregates (Chemical Method)<sup>2</sup>
- C 330 Specification for Lightweight Aggregates for Structural Concrete<sup>2</sup>
- C 331 Specification for Lightweight Aggregates for Concrete Masonry Units<sup>2</sup>
- C 332 Specification for Lightweight Aggregates for Insulating Concrete<sup>2</sup>
- C 638 Constituents of Aggregates for Radiation-Shielding Concrete<sup>2</sup>

<sup>1</sup> This descriptive nomenclature is under the jurisdiction of ASTM Committee C09 on Concrete and Concrete Aggregates and is the direct responsibility of Subcommittee C09.65 on Petrography.

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This standard has been extensively revised. The reader should compare this edition with the last previous edition for exact revisions.

<sup>2</sup> Annual Book of ASTM Standards, Vol 04.02.

### 3. Significance and Use

3.1 This descriptive nomenclature provides information on terms commonly applied to concrete aggregates. This standard is intended to assist in understanding the meaning and significance of the terms.

3.2 Many of the materials described frequently occur in particles that do not display all the characteristics given in the descriptions, and most of the described rocks grade from varieties meeting one description to varieties meeting another with all intermediate stages being found.

3.3 The accurate identification of rocks and minerals can, in many cases, be made only by a qualified geologist, mineralogist, or petrographer using the apparatus and procedures of these sciences. Reference to these descriptions may, however, serve to indicate or prevent gross errors in identification. Identification of the constituent materials in an aggregate may assist in characterizing its engineering properties, but identification alone cannot provide the sole basis for predicting behavior of aggregates in service. Aggregates of any type or combination of types may perform well or poorly in service depending upon the exposure to which the concrete is subjected, the physical and chemical properties of the matrix in which they are embedded, their physical condition at the time they are used, and other factors. Constituents that may occur only in minor amounts in the aggregate may or may not decisively influence its performance. Information about concrete aggregate performance in concrete has been published by ASTM.<sup>3</sup>

## CONSTITUENTS OF NATURAL MINERAL AGGREGATES

### 4. Classes and Types

4.1 The materials found as constituents of natural mineral aggregates are minerals and rocks.

4.2 Minerals are naturally occurring inorganic substances of more or less definite chemical composition and usually of a specific crystalline structure. Most rocks are composed of several minerals but some are composed of only one mineral.

<sup>3</sup> Klieger, P., and Lamond, J. F., editors, Significance of Tests and Properties of Concrete and Concrete-Making Materials, ASTM STP 169C, 1994.

Certain examples of the rock quartzite are composed exclusively of the mineral quartz, and certain limestones are composed exclusively of the mineral calcite. Individual sand grains frequently are composed of particles of rock, but they may be composed of a single mineral, particularly in the finer sizes.

4.3 Rocks are classified according to origin into three major divisions: igneous, sedimentary, and metamorphic. These three major groups are subdivided into types according to mineral and chemical composition, texture, and internal structure.

4.3.1 Igneous rocks form from molten matter either at or below the earth's surface.

4.3.2 Sedimentary rocks form near the earth's surface by the accumulation and consolidation of the products of weathering and erosion of existing rocks, or by direct chemical precipitation. Sedimentary rocks may form from pre-existing igneous, metamorphic, or sedimentary rocks.

4.3.3 Metamorphic rocks form from pre-existing igneous, sedimentary, or metamorphic rocks by the action of heat or pressure or both.

## 5. Silica Minerals

5.1 *Quartz*—a very common hard mineral composed of silica ( $\text{SiO}_2$ ). It will scratch glass and is not scratched by a knife. When pure it is colorless with a glassy (vitreous) luster and a shell-like (conchoidal) fracture. It lacks a visible cleavage (the ability to break in definite directions along even planes) and, when present in massive rocks such as granite, it usually has no characteristic shape. It is resistant to weathering and is therefore an important constituent of many sand and gravel deposits and many sandstones. It is also abundant in many light-colored igneous and metamorphic rocks. Some strained, or intensely fractured (granulated), and microcrystalline quartz may be potentially deleteriously reactive with the alkalis in the hydraulic cement paste.

5.2 *Opal*—a hydrous form of silica ( $\text{SiO}_2 \cdot n\text{H}_2\text{O}$ ) which occurs without characteristic external form or internal crystalline arrangement as determined by ordinary visible light methods. When X-ray diffraction methods are used, opal may show some evidences of internal crystalline arrangement. Opal has a variable water content, generally ranging from 3 to 9%. The specific gravity and hardness are always less than those of quartz. The color is variable and the luster is resinous to glassy. It is usually found in sedimentary rocks, especially some cherts, and is the principal constituent of diatomite. It is also found as a secondary material filling cavities and fissures in igneous rocks and may occur as a coating on gravel and sand. The recognition of opal in aggregates is important because it is potentially deleteriously reactive with the alkalis in hydraulic cement paste or with the alkalis from other sources, such as aggregates containing zeolites, and ground water.

5.3 *Chalcedony*—chalcedony has been considered both as a distinct mineral and a variety of quartz. It is frequently composed of a mixture of microscopic fibers of quartz with a large number of submicroscopic pores filled with water and air. The properties of chalcedony are intermediate between those of opal and quartz, from which it can sometimes be distinguished only by laboratory tests. It frequently occurs as a constituent of

the rock chert and is potentially deleteriously reactive with the alkalis in hydraulic cement paste.

5.4 *Tridymite and cristobalite*—high temperature crystalline forms of silica ( $\text{SiO}_2$ ) sometimes found in volcanic rocks. They are metastable at ordinary temperatures and pressures. They are rare minerals in aggregates except in areas where volcanic rocks are abundant. A type of cristobalite is a common constituent of opal. Tridymite and cristobalite are potentially deleteriously reactive with the alkalis in hydraulic cement paste.

## 6. Feldspars

6.1 The minerals of the feldspar group are the most abundant rock-forming minerals in the crust of the earth. They are important constituents of all three major rock groups, igneous, sedimentary, and metamorphic. Since all feldspars have good cleavages in two directions, particles of feldspar usually show several smooth surfaces. Frequently, the smooth cleavage surfaces show fine parallel lines. All feldspars are slightly less hard than, and can be scratched by, quartz and will, when fresh, easily scratch a penny. The various members of the group are differentiated by chemical composition and crystallographic properties. The feldspars *orthoclase*, *sanidine*, and *microcline* are potassium aluminum silicates, and are frequently referred to as *potassium feldspars*. The *plagioclase* feldspars include those that are sodium aluminum silicates and calcium aluminum silicates, or both sodium and calcium aluminum silicates. This group, frequently referred to as the “soda-lime” group, includes a continuous series, of varying chemical composition and optical properties, from *albite*, the sodium aluminum feldspar, to *anorthite*, the calcium aluminum feldspar, with intermediate members of the series designated *oligoclase*, *andesine*, *labradorite*, and *bytownite*. Potassium feldspars and sodium-rich plagioclase feldspars occur typically in igneous rocks such as granites and rhyolites, whereas, plagioclase feldspars of higher calcium content are found in igneous rocks of lower silica content such as diorite, gabbro, andesite, and basalt.

## 7. Ferromagnesian Minerals

7.1 Many igneous and metamorphic rocks contain dark green to black minerals that are generally silicates of iron or magnesium, or of both. They include the minerals of the amphibole, pyroxene, and olivine groups. The most common amphibole mineral is hornblende; the most common pyroxene mineral is augite; and the most common olivine mineral is forsterite. Dark mica, such as biotite and phlogopite, are also considered ferromagnesian minerals. The amphibole and pyroxene minerals are brown to green to black and generally occur as prismatic units. Olivine is usually olive green, glassy in appearance, and usually altered. Biotite has excellent cleavage and can be easily cleaved into thin flakes and plates. These minerals can be found as components of a variety of rocks, and in sands and gravels. Olivine is found only in dark igneous rocks where quartz is not present, and in sands and gravels close to the olivine source.

## 8. Micaceous Minerals

8.1 Micaceous minerals have perfect cleavage in one direction and can be easily split into thin flakes. The mica minerals of the muscovite group are colorless to light green; of the biotite group, dark brown to black or dark green; of the lepidolite group, white to pink and red or yellow; and of the chlorite group, shades of green. Another mica, phlogopite, is similar to biotite, commonly has a pearl-like luster and bronze color, and less commonly is brownish red, green, or yellow. The mica minerals are common and occur in igneous, sedimentary, and metamorphic rocks, and are common as minor to trace components in many sands and gravels. The muscovite, biotite, lepidolite, and phlogopite minerals cleave into flakes and plates that are elastic; the chlorite minerals, by comparison, form in elastic flakes and plates. Vermiculite (a mica-like mineral) forms by the alteration of other micas and is brown and has a bronze luster.

## 9. Clay Minerals

9.1 The term “clay” refers to natural material composed of particles in a specific size range less than 2  $\mu\text{m}$  (0.002 mm). Mineralogically, clay refers to a group of layered silicate minerals including the clay-micas (illites), the kaolin group, very finely divided chlorites, and the swelling clays—smectite including montmorillonites. Members of several groups, particularly micas, chlorites, and vermiculites, occur both in the clay-size range and in larger sizes. Some clays are made up of alternating layers of two or more clay groups. Random, regular, or both types of interlayering are known. If smectite is a significant constituent in such mixtures, then fairly large volume changes may occur with wetting and drying.

9.2 Clay minerals are hydrous aluminum, magnesium, and iron silicates that may contain calcium, magnesium, potassium, sodium, and other exchangeable cations. They are formed by alteration and weathering of other silicates and volcanic glass. The clay minerals are major constituents of clays and shales. They are found disseminated in carbonate rocks as seams and pockets and in altered and weathered igneous and metamorphic rocks. Clays may also be found as matrix, void fillings, and cementing material in sandstones and other sedimentary rocks.

9.3 Most aggregate particles composed of, or containing, large proportions of clay minerals are soft and, because of the large internal surface area of the constituents, they are porous. Some of these aggregates will disintegrate when wetted. Rocks in which the cementing matrix is principally clay, such as clay-bonded sandstones, and rocks in which swelling clay minerals (smectite) are present as a continuous phase or matrix, such as in some altered volcanics, may slake in water or may disintegrate in the concrete mixer. Rocks of this type are unsuitable for use as aggregates. Rocks having these properties less well developed will abrade considerably during mixing, releasing clay, and raising the water requirement of the concrete containing them. When such rocks are present in hardened concrete, the concrete will manifest greater volume change on wetting and drying than similar concrete containing non-swelling aggregate.

## 10. Zeolites

10.1 The zeolite minerals are a large group of hydrated aluminum silicates of the alkali and alkaline earth elements which are soft and usually white or light colored. They are formed as a secondary filling in cavities or fissures in igneous rocks, or within the rock itself as a product of hydrothermal alteration of original minerals, especially feldspars. Some zeolites, particularly *heulandite*, *natrolite*, and *laumontite*, reportedly produce deleterious effects in concrete, the first two having been reported to raise the alkali content in concrete by releasing alkalis through cation exchange and thus increasing alkali reactivity when alkali-reactive aggregate constituents are present. Laumontite and its partially dehydrated variety *leonhardite* are notable for their substantial volume change with wetting and drying. Both are found in rocks such as quartz diorites and some sandstones.

## 11. Carbonate Minerals

11.1 The most common carbonate mineral is *calcite* (calcium carbonate,  $\text{CaCO}_3$ ). The mineral *dolomite* consists of calcium carbonate and magnesium carbonate ( $\text{CaCO}_3 \cdot \text{MgCO}_3$  or  $\text{CaMg}(\text{CO}_3)_2$ ) in equivalent molecular amounts, which are 54.27 and 45.73 by mass %, respectively. Both calcite and dolomite are relatively soft, the hardness of calcite being 3 and that of dolomite  $3\frac{1}{2}$  to 4 on the Mohs scale, and are readily scratched by a knife blade. They have rhombohedral cleavage, which results in their breaking into fragments with smooth parallelogram shaped sides. Calcite is soluble with vigorous effervescence in cold dilute hydrochloric acid; dolomite is soluble with slow effervescence in cold dilute hydrochloric acid and with vigorous effervescence if the acid or the sample is heated or if the sample is pulverized.

## 12. Sulfate Minerals

12.1 Carbonate rocks and shales may contain sulfates as impurities. The most abundant sulfate mineral is *gypsum* (hydrous calcium sulfate;  $\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$ ); *anhydrite* (anhydrous calcium sulfate,  $\text{CaSO}_4$ ) is less common. Gypsum is usually white or colorless and characterized by a perfect cleavage along one plane and by its softness, representing hardness of 2 on the Mohs scale; it is readily scratched by the fingernail. Gypsum may form a whitish pulverulent or crystalline coating on sand and gravel. It is slightly soluble in water.

12.2 Anhydrite resembles dolomite in hand specimen but has three cleavages at right angles; it is less soluble in hydrochloric acid than dolomite, does not effervesce and is slightly soluble in water. Anhydrite is harder than gypsum. Gypsum and anhydrite occurring in aggregates can cause sulfate attack in concrete and mortar.

## 13. Iron Sulfide Minerals

13.1 The sulfides of iron, *pyrite*, *marcasite*, and *pyrrhotite* are frequently found in natural aggregates. Pyrite is found in igneous, sedimentary, and metamorphic rocks; marcasite is much less common and is found mainly in sedimentary rocks; pyrrhotite is less common but may be found in many types of igneous and metamorphic rocks. Pyrite is brass yellow, and pyrrhotite bronze brown, and both have a metallic luster.

Marcasite is also metallic but lighter in color and finely divided iron sulfides are soot black. Pyrite is often found in cubic crystals. Marcasite readily oxidizes with the liberation of sulfuric acid and formation of iron oxides, hydroxides, and, to a much smaller extent, sulfates; pyrite and pyrrhotite do so less readily. Marcasite and certain forms of pyrite and pyrrhotite are reactive in mortar and concrete, producing a brown stain accompanied by a volume increase that has been reported as one source of popouts in concrete. Reactive forms of iron sulfides may be recognized by immersion in saturated lime water (calcium hydroxide solution); upon exposure to air the reactive varieties produce a brown coating within a few minutes.

#### 14. Iron Oxide Minerals, Anhydrous and Hydrous

14.1 There are two common iron oxide minerals: (1) Black, magnetic: *magnetite* ( $\text{Fe}_3\text{O}_4$ ), and (2) red or reddish when powdered: *hematite* ( $\text{Fe}_2\text{O}_3$ ); and one common hydrous oxide mineral, brown or yellowish: *goethite* ( $\text{FeO}(\text{OH})$ ). Another common iron-bearing mineral is black, weakly magnetic, *ilmenite* ( $\text{FeTiO}_3$ ). Magnetite and ilmenite are important accessory minerals in many dark igneous rocks and are common detrital minerals in sediments. Hematite is frequently found as an accessory mineral in reddish rocks. Limonite, the brown weathering product of iron-bearing minerals, is a field name for several varieties of hydrous iron oxide minerals including goethite; it frequently contains adsorbed water, and various impurities such as colloidal or crystalline silica, clay minerals, and organic matter. The presence of substantial amounts of soft iron-oxide minerals in concrete aggregate can color concrete various shades of yellow or brown. Very minor amounts of iron minerals color many rocks, such as ferruginous sandstones, shales, clay-ironstones, and granites. Magnetite, ilmenite, and hematite ores are used as heavy aggregates.

### DESCRIPTIONS OF IGNEOUS ROCKS

#### 15. General

15.1 Igneous rocks are those formed by cooling from a molten rock mass (magma). They may be divided into two classes: (1) plutonic, or intrusive, that have cooled slowly within the earth; and (2) volcanic, or extrusive, that formed from quickly cooled lavas. Plutonic rocks have grain sizes greater than approximately 1 mm, and are classified as coarse- or medium-grained. Volcanic rocks have grain sizes less than approximately 1 mm, and are classified as fine-grained. Volcanic rocks frequently contain glass. Both plutonic and volcanic rocks may consist of porphyries, that are characterized by the presence of large mineral grains in a fine-grained or glassy groundmass. This is the result of sharp changes in rate of cooling or other physico-chemical conditions during solidification of the melt.

15.2 Igneous rocks are usually classified and named on the basis of their texture, internal structure, and their mineral composition which in turn depends to a large extent on their chemical composition. Rocks in the plutonic class generally have chemical equivalents in the volcanic class.

#### 16. Plutonic Rocks

16.1 *Granite*—granite is a medium- to coarse-grained, light-colored rock characterized by the presence of potassium feldspar with lesser amounts of plagioclase feldspars and quartz. The characteristic potassium feldspars are orthoclase or microcline, or both; the common plagioclase feldspars are albite and oligoclase. Feldspars are more abundant than quartz. Dark-colored mica (biotite) is usually present, and light-colored mica (muscovite) is frequently present. Other dark-colored ferromagnesian minerals, especially hornblende, may be present in amounts less than those of the light-colored constituents. *Quartz-monzonite* and *granodiorite* are rocks similar to granite, but they contain more plagioclase feldspar than potassium feldspar.

16.2 *Syenite*—syenite is a medium- to coarse-grained, light-colored rock composed essentially of alkali feldspars, namely microcline, orthoclase, or albite. Quartz is generally absent. Dark ferromagnesian minerals such as hornblende, biotite, or pyroxene are usually present.

16.3 *Diorite*—diorite is a medium- to coarse-grained rock composed essentially of plagioclase feldspar and one or more ferromagnesian minerals such as hornblende, biotite, or pyroxene. The plagioclase is intermediate in composition, usually of the variety andesine, and is more abundant than the ferromagnesian minerals. Diorite usually is darker in color than granite or syenite and lighter than gabbro. If quartz is present, the rock is called *quartz diorite*.

16.4 *Gabbro*—gabbro is a medium- to coarse-grained, dark-colored rock consisting essentially of ferromagnesian minerals and plagioclase feldspar. The ferromagnesian minerals may be pyroxenes, amphiboles, or both. The plagioclase is one of the calcium-rich varieties, namely labradorite, bytownite, or anorthite. Ferromagnesian minerals are usually more abundant than feldspar. *Diabase* (in European usage *dolerite*) is a rock of similar composition to gabbro and basalt but is intermediate in mode of origin, usually occurring in smaller intrusions than gabbro, and having a medium to fine-grained texture. The terms “trap” or “trap rock” are collective terms for dark-colored, medium- to fine-grained igneous rocks especially diabase and basalt.

16.5 *Peridotite*—peridotite is composed of olivine and pyroxene. Rocks composed almost entirely of pyroxene are known as *pyroxenites*, and those composed of olivine as *dunites*. Rocks of these types are relatively rare but their metamorphosed equivalent, serpentinite, is more common.

16.6 *Pegmatite*—extremely coarse-grained varieties of igneous rocks are known as pegmatites. These are *usually* light-colored and are most frequently equivalent to granite or syenite in mineral composition.

#### 17. Fine-Grained and Glassy Extrusive Igneous Rocks

17.1 *Volcanic Rock*—volcanic or extrusive rocks are the fine-grained equivalents of the coarse-and-medium-grained plutonic rocks. Equivalent types have similar chemical compositions and may contain the same minerals. Volcanic rocks commonly are so fine-grained that the individual mineral grains usually are not visible to the naked eye. Porphyritic textures are common, and the rocks may be partially or wholly

glassy or non-crystalline. The glassy portion of a partially glassy rock usually has a higher silica content than the crystalline portion. Some volcanic or extrusive rocks may not be distinguishable in texture and structure from plutonic or intrusive rocks that originated at shallow depth.

17.2 *Glassy Volcanic Rocks*—These rocks are of particular significance because they contain, or may contain, high silica glass that is alkali-reactive, and secondary minerals that are alkali-reactive or release alkalies into concrete. The high silica glasses, generally classed as those containing more than 55 % silica, are known to be alkali-reactive, whereas the low-silica glasses are not. Among igneous rocks that contain, or may contain, high silica glass are: obsidian, pumice, trachyte, rhyolite, scoria, dacite, basalt, andesite, and perlite. Aggregates containing these rocks include crushed parent rock where the aggregate is constituted totally of the rock, or as varying percentages in gravels and sands. Glassy rocks, particularly the more siliceous ones, are potentially deleteriously reactive with the alkalies in hydraulic cement paste.

17.3 *Volcanic Glass*—igneous rocks composed wholly of glass are named on the basis of their texture and internal structure. A dense dark natural glass of high silica content is called *obsidian*, while lighter colored finely vesicular glassy froth filled with elongated, tubular bubbles is called *pumice*. Dark-colored coarsely vesicular types containing more or less spherical bubbles are called *scoria*. Pumices are usually silica-rich (corresponding to rhyolites or dacites), whereas scorias usually are more basic (corresponding to basalts). A high-silica glassy lava with an onion-like structure and a pearly luster, containing 2 to 5 % water, is called *perlite*. When heated quickly to the softening temperature, perlite puffs to become an artificial pumice. Glass with up to 10 % water and with a dull resinous luster is called *pitchstone*.

17.4 *Felsite*—light-colored, very fine-grained igneous rocks are collectively known as felsites. The felsite group includes *rhyolite*, *dacite*, *andesite*, and *trachyte*, which are the equivalents of granite, quartz diorite, diorite, and syenite, respectively. These rocks are usually light colored but they may be gray, green, dark red, or black. When they are microcrystalline or contain natural glass, rhyolites, dacites, and andesites are potentially deleteriously reactive with the alkalies in hydraulic cement paste.

17.5 *Basalt*—fine-grained extrusive equivalent of gabbro and diabase. When basalt contains natural glass, the glass is generally lower in silica content than that of the lighter-colored extrusive rocks and hence is not deleteriously reactive with the alkalies in hydraulic cement paste; however, exceptions have been noted in the literature with respect to the alkali reactivity of basaltic glasses.

17.6 Vesicles and other voids in volcanic rocks may contain alkali-reactive forms of silica such as opal, cristobalite, tridymite, and various varieties of microcrystalline quartz. Additionally, zeolitic minerals can release alkalies and thus increase the alkalies in the paste.

## DESCRIPTIONS OF SEDIMENTARY ROCKS

### 18. General

18.1 Sedimentary rocks are stratified rocks usually laid down under water, although they can also be formed by wind and glacial action. Sediments may be composed of particles of preexisting rocks derived by mechanical agencies or they may be of chemical or organic origin. The sediments are usually indurated by cementation or compaction during geologic time, although the degree of consolidation may vary widely.

18.2 *Gravel, sand, silt, and clay* form the group of unconsolidated sediments. Although the distinction between these four members is made on the basis of their particle size, a general trend in the composition occurs. Gravel and, to a lesser degree, coarse sands usually consist of rock fragments; fine sands and silt consist predominantly of mineral grains; and clay exclusively of mineral grains, largely of the group of clay minerals. All types of rocks and minerals may be represented in unconsolidated sediments.

### 19. Conglomerates, Sandstones, and Quartzites

19.1 These rocks consist of particles of *sand* or *gravel*, or both, with or without interstitial and cementing material. If the particles include a considerable proportion of gravel, the rock is a *conglomerate*. If the particles are in the sand sizes, that is, less than 2 mm but more than 0.06 mm in major diameter, the rock is a *sandstone* or a *quartzite*. If the rock breaks around the sand grains, it is a sandstone; if the grains are largely quartz and the rock breaks through the grains, it is quartzite. Conglomerates, and sandstones are sedimentary rocks but quartzites may be sedimentary (*orthoquartzites*) or metamorphic (*metaquartzites*). The cementing or interstitial materials of sandstones may be quartz, opal, calcite, dolomite, clay, iron oxides, or other materials. These may influence the quality of a sandstone as concrete aggregate. If the nature of the cementing material is known, the rock name may include a reference to it, such as *opal-bonded sandstone* or *ferruginous conglomerate*. Opal-containing rocks may be potentially deleteriously reactive with alkalies in the hydraulic cement paste.

19.2 *Graywackes and subgraywackes*—gray to greenish gray sandstones containing angular quartz and feldspar grains, and sand-sized rock fragments in an abundant matrix resembling claystone, shale, argillite, or slate. Graywackes grade into subgraywackes, the most common sandstones of the geologic column.

19.3 *Arkose*—coarse-grained sandstone derived from granite, containing conspicuous amounts of feldspar.

### 20. Claystones, Shales, Argillites, and Siltstones

20.1 These very fine-grained rocks are composed of, or derived by erosion of sedimentary *silts* and *clays*, or of any type of rock that contained clay. When relatively soft and massive, they are known as *claystones*, or *siltstones*, depending on the size of the majority of the particles of which they are composed. Siltstones consist predominantly of silt-sized particles (0.0625 to 0.002 mm in diameter) and are intermediate

rocks between claystones and sandstones. When the claystones are harder and platy or fissile, they are known as *shales*. Claystones and shales may be gray, black, reddish, or green and may contain some carbonate minerals (calcareous shales). A massive, firmly indurated fine-grained argillaceous rock consisting of quartz, feldspar, and micaceous minerals is known as *argillite*. Argillites do not slake in water as some shales do. As an aid in distinguishing these fine-grained sediments from fine-grained, foliated metamorphic rocks such as slates and phyllites, it may be noted that the cleavage surfaces of shales are generally dull and earthy while those of slates are more lustrous. Phyllite has a glossier luster resembling a silky sheen.

20.2 Aggregates containing abundant shale may be detrimental to concrete because they can produce high shrinkage, but not all shales are harmful. Some argillites are alkali-silica reactive.

20.3 Although aggregates which are volumetrically unstable in wetting and drying are not confined to any class of rock, they do share some common characteristics. If there is a matrix or continuous phase, it is usually physically weak and consists of material of high specific surface, frequently including clay. However, no general relation has been demonstrated between clay content or type of clay and large volume change upon wetting and drying. Volumetrically unstable aggregates do not have mineral grains of high modulus interlocked in a continuous rigid structure capable of resisting volume change.

20.4 Aggregates having high elastic modulus and low volume change from the wet to the dry condition contribute to the volume stability of concrete by restraining the volume change of the cement paste. In a relatively few cases, aggregates have been demonstrated to contribute to unsatisfactory performance of concrete because they have relatively large volume change from the wet to the dry condition combined with relatively low modulus of elasticity. On drying, such aggregates shrink away from the surrounding cement paste and consequently fail to restrain its volume change with change in moisture content.

## 21. Carbonate Rocks

21.1 *Limestones*, the most widespread of carbonate rocks. They range from pure limestones consisting of the mineral calcite to pure *dolomites* (dolostones) consisting of the mineral dolomite. Usually they contain both minerals in various proportions. If 50 to 90 % is the mineral dolomite, the rock is called *calcitic dolomite*. The term “magnesium limestone” is sometimes applied to dolomitic limestones and calcitic dolomites but it is ambiguous and its use should be avoided. Most carbonate rocks contain some noncarbonate impurities such as quartz, chert, clay minerals, organic matter, gypsum, and sulfides. Carbonate rocks containing 10 to 50 % sand are *arenaceous* (or *sandy*) *limestones* (or *dolomites*); those containing 10 to 50 % clay are *argillaceous* (or *clayey* or *shaly*) *limestones* (or *dolomites*). *Marl* is a clayey limestone which is fine-grained and commonly soft. Chalk is fine-textured, very soft, porous, and somewhat friable limestone, composed chiefly of particles of microorganisms. Micrite is very fine-textured chemically precipitated carbonate or a mechanical ooze of carbonate particles, usually 0.001 to 0.003 mm in size. The term “limerock” is not recommended.

21.2 The reaction of the dolomite in certain carbonate rocks with alkalis in portland cement paste has been found to be associated with deleterious expansion of concrete containing such rocks as coarse aggregate. Carbonate rocks capable of such reaction possess a characteristic texture and composition. The characteristic microscopic texture is that in which relatively large crystals of dolomite (rhombs) are scattered in a finer-grained matrix of micritic calcite and clay. The characteristic composition is that in which the carbonate portion consists of substantial amounts of both dolomite and calcite, and the acid-insoluble residue contains a significant amount of clay. Except in certain areas, such rocks are of relatively infrequent occurrence and seldom make up a significant proportion of the material present in a deposit of rock being considered for use in making aggregate for concrete.

## 22. Chert

22.1 *Chert*—the general term for a group of variously colored, very fine-grained (aphanitic), siliceous rocks composed of microcrystalline or cryptocrystalline quartz, chalcedony, or opal, either singly or in combinations of varying proportions. Identification of the form or forms of silica requires careful determination of optical properties, absolute specific gravity, loss on ignition, or a combination of these characteristics. Dense cherts are very tough, with a waxy to greasy luster, and are usually gray, brown, white, or red, and less frequently, green, black or blue. Porous varieties are usually lighter in color, frequently off-white, or stained yellowish, brownish, or reddish, firm to very weak, and grade to tripoli. Ferruginous, dense, red, and in some cases, dense, yellow, brown, or green chert is sometimes called *jasper*. Dense black or gray chert is sometimes called *flint*. A very dense, even textured, light gray to white chert, composed mostly of microcrystalline to cryptocrystalline quartz, is called *novaculite*. Chert is hard (scratches glass, but is not scratched by a knife blade) and has a conchoidal (shell-like) fracture in the dense varieties, and a more splintery fracture in the porous varieties. Chert occurs most frequently as nodules, lenses, or interstitial material, in limestone and dolomite formations, as extensively bedded deposits, and as components of sand and gravel. Most cherts have been found to be alkali-silica reactive to some degree when tested with high-alkali cement, or in the quick chemical test (Test Method C 289). In the absence of information to the contrary, all chert should be regarded as potentially deleteriously reactive with the alkalis in hydraulic cement paste. The degree of alkali-silica reactivity, and whether a given chert will produce a deleterious degree of expansion in concrete, are complex functions of several factors. The degree of the alkali-silica reactivity and whether a given chert will produce a deleterious degree of expansion in concrete are complex functions of several factors. Among them are: the mineralogic composition and internal structure of the chert; the amount of the chert as a proportion of the aggregates; the particle-size distribution; the alkali content of the cement; and the cement content of the concrete. However, opaline cherts may produce deleterious expansion of mortar or concrete when present in very small proportions (less than 5 % by mass of the aggregate). Cherts that are porous may be susceptible to freezing and thawing deterioration in concrete